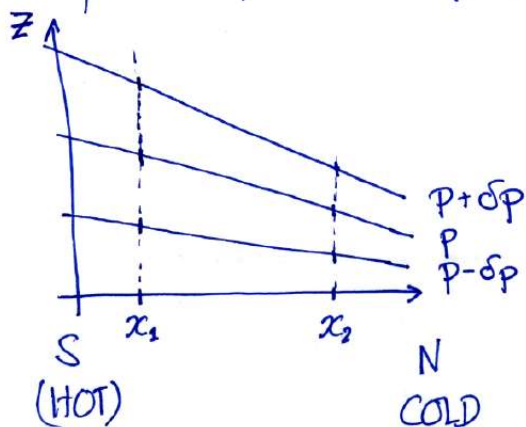


Basics of thermal wind

1/

- ① geostrophic relationship says geostrophic winds are driven by horizontal gradients in the geopotential.
- ② secondly, geopotential thickness is a function of the mean temperature gradient in the layer
- ③ so horizontal gradients in temperature drive horizontal gradients in thickness
- ④ these horizontal gradients in thickness imply that the horizontal gradients in geopotential vary as a function of height. Therefore, because of ①, geostrophic winds vary with height.

Consider pressure surfaces as a fn of height & latitude ^{z,}



from geopotential equation $\delta z = \frac{RT}{g_0} \delta \ln p$ where \bar{T} = mean temperature within layer $\delta \ln p$

- because it is hot in the south, the thickness between pressure surfaces is greater.
- therefore, because of the increasing tilt of constant pressure surfaces as z increases, the gradient $(\partial \phi / \partial x)_p$ increases.
- therefore the geostrophic wind must increase.

Maths of thermal wind

Geostrophic balance $u = -\frac{1}{f} \frac{d\phi}{dy}_p$

$$d/dp : \Rightarrow \frac{du}{dp} = -\frac{1}{f} \frac{d}{dy} \frac{d\phi}{dp}$$

now $dp = -g\rho dz = -\rho d\phi = -\frac{\rho d\phi}{RT}$

$$\Rightarrow \frac{d\phi}{dp} = -RT/p$$

substitute $\frac{du}{dp} = +\frac{R}{f p} \frac{dT}{dy}_p$

[have related horizontal temperature gradient to vars of u with height]

$$\Rightarrow p \frac{du}{dp} = \frac{du}{d \ln p} = \frac{R}{f} \frac{dT}{dy}_p$$

now, $p \frac{d}{dp} = -\frac{RT}{g} \frac{d}{dz}$

[nb note the minus sign]

⇒

$$\frac{dv}{dz} = -\frac{g}{fT} \frac{dT}{dy}_z$$

[nb making the approximation that $\frac{dT}{dy}_z = \frac{dT}{dy}_p$]

or in vector form

$$\frac{d \underline{v}_h}{dz} = \frac{g}{fT} \underline{k} \times \nabla T$$

- ⇒ therefore in the presence of a horizontal ~~wind~~ temperature gradient, the geostrophic wind increases with height.
- ⇒ The direction of that increase is PERPENDICULAR to the temperature gradient.